

“Convection Resolving Model” (CRM) MOLOCH



1-Breve descrizione del CRM sviluppato all'ISAC-CNR

2-Ipotesi alla base della parametrizzazione dei processi
microfisici



Objectives

- Develop a tool for very high resolution-short range operational weather forecast and Nowcasting;
- ‘Resolve’ explicitly atmospheric convection (without parameterization);
- Develop a tool for research purposes (simulation of thunderstorm development, flows over complex orography, physical processes responsible for intense precipitation,

Model dynamics

- non hydrostatic, fully compressible;
- Arakawa C grid; terrain-following coordinate
- time split, implicit for vertically propagating sound waves, FB for horizontal prop. waves
- advection: FBAS (Malguzzi & Tartaglione, 1999); also Weighted Average Flux WAF (Toro 1989; Hubbard & Nikiforakis, 2001)
- nested in BOLAM runs

Model physics

- radiation, vertical diffusion, surface turbulent fluxes similar to BOLAM
- soil water and energy balance based on Pressman soil scheme
- cloud microphysics (partly based on Drofa, 2003)
- no dry and moist convection

Governing equations

$$\frac{du}{dt} = -\frac{1}{\rho} \frac{\partial P}{\partial x} - fv + K_u$$

$$\frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial P}{\partial y} + fu + K_v$$

$$\frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial P}{\partial z} - g$$

$$\frac{dT}{dt} = -T \frac{R'}{C_V} \vec{\nabla} \cdot \vec{V} + \frac{\dot{Q}}{C_V} + K_T$$

$$\frac{dP}{dt} = -P \gamma \vec{\nabla} \cdot \vec{V} + \frac{P \dot{Q}}{C_V T}$$

$$P = \rho R' T$$

Effective gas constant

$$R' = R_d \left(1 + (1/\varepsilon - 1) q_V - q_W - q_I \right)$$

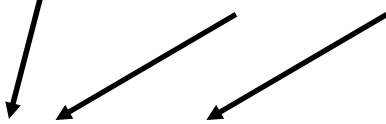
$$\gamma = \frac{C_P}{C_V} \quad \varepsilon = \frac{R_d}{R_V}$$

Governing equations:
Conservation of specific concentration of
water species in air parcels

$$\frac{dq_k}{dt} = \delta_{kV} K_V + \dots\dots\dots$$

$k = V, Cw, Ci, Pw, Pi_1, Pi_2$

Hydrometeors



Turbulent kinetic energy equation in H -coordinates

$$\frac{d\bar{E}}{dt} = K_E - \overline{u_i' u_j'^i} \frac{\partial \bar{u}_i}{\partial x_j} \Big|_s + \frac{g}{\bar{\theta}_V} \overline{w' \theta'_V} - \varepsilon, \quad \vec{u}' = \begin{pmatrix} u' \\ v' \\ w' \end{pmatrix}, \quad \vec{u}'^i = \begin{pmatrix} u' \\ v' \\ s' \end{pmatrix}$$

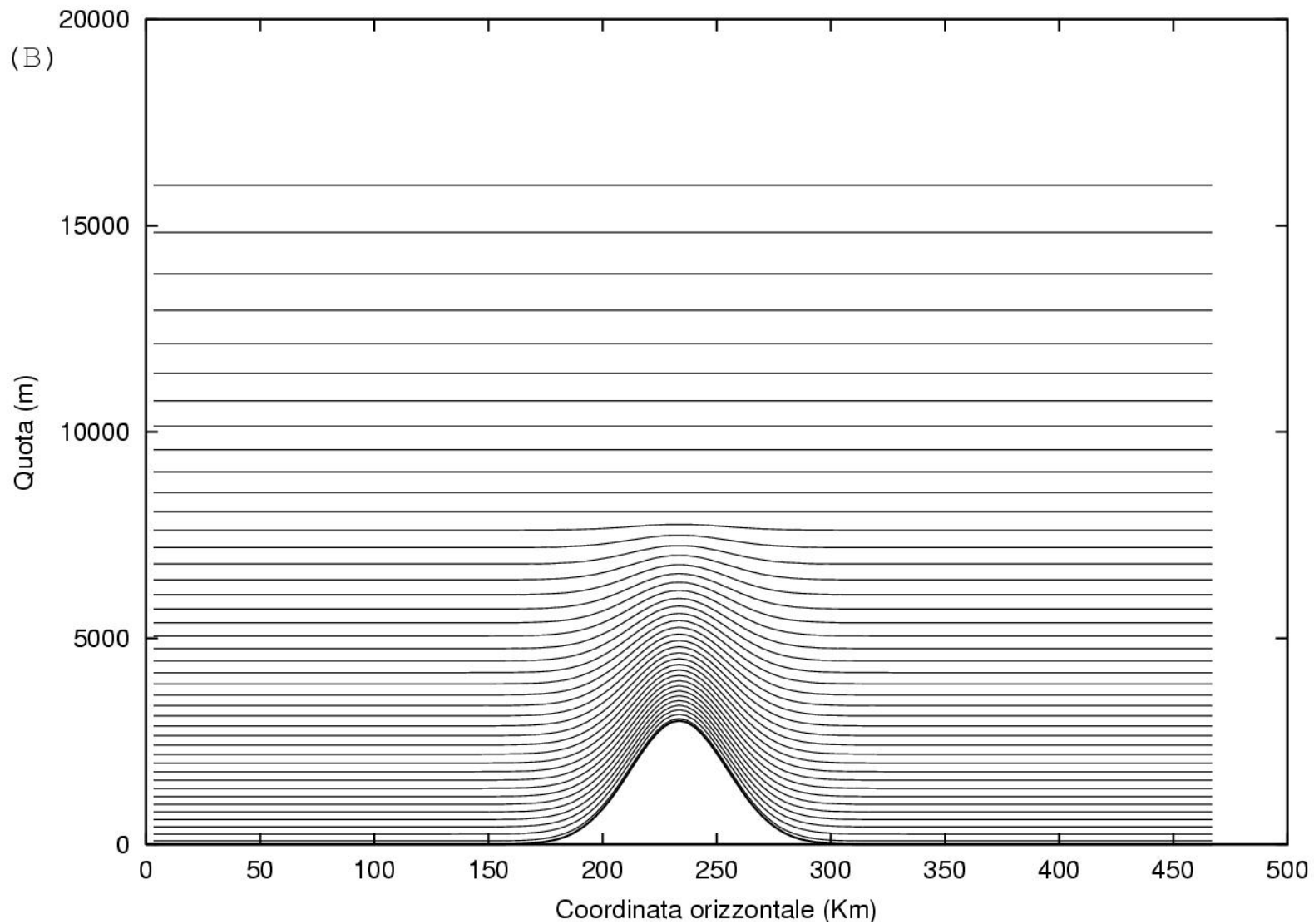
$$\bar{E} = \frac{1}{2} (\overline{u'u'} + \overline{v'v'} + \overline{w'w'})$$

$$\varepsilon = \frac{1}{l_m} \left(\sqrt{C_E \bar{E}} \right)^3$$

Closure: Energy redistribution hypothesis

$$\overline{u_i' u_j'} = \frac{2}{3} \delta_{ij} E - K \left(\frac{\partial \bar{u}_i}{\partial x_j} + \frac{\partial \bar{u}_j}{\partial x_i} \right) \quad K = l_m \sqrt{C_E E}$$

Hybrid coordinates ' H '



H -Coordinate

Terrain following vertical coordinate

$$\zeta = H \left(1 - e^{-\frac{z-h(1-\zeta/H)}{H}} \right)$$

$$h < z < \infty$$

The vertical scale H is given by the *density scale height*

$$H = \frac{R_d T_0}{g}$$

Microphysical hypothesis

Liquid and solid cloud particles:
gamma-distribution (Levi distribution) for
the number of particles per unit volume
and unit radius D :

$$N(D) = \frac{N_0 \beta^{\alpha+1}}{\Gamma(\alpha+1)} D^\alpha e^{-\beta D}$$

$N_0 = 8 \cdot 10^6 \text{ (m}^{-3}\text{)}$ and $\alpha = 6$ for cloud drops
 $N_0 = 2 \cdot 10^7 \text{ (m}^{-3}\text{)}$ and $\alpha = 3$ for cloud crystal

Liquid and solid precipitation:
Marshall-Palmer distribution

$$N(D) = N_0 e^{-\lambda D}$$

$N_0 = 8 \cdot 10^6 \text{ (m}^{-4}\text{)}$ for precipitating drops
 N_0 function of crystal shape for precipitating ice

β and λ are determined from the normalizing condition: $\frac{1}{\rho} \int N(D) m(D) dD = q$

where m is the mass of a particle of diameter D : $m(D) = aD^b$

and where $a = \pi/6 \cdot \rho_w$, $b = 3$ for cloud and precipitating water; $a = 100$, $b = 2.5$ for cloud ice;
 a and b function of crystal shape (temperature) for precipitating ice. The result is:

$$\beta = \left[\frac{N_0 a \Gamma(\alpha + b + 1)}{\rho q \Gamma(\alpha + 1)} \right]^{\frac{1}{b}}$$

$$\lambda = \left[\frac{N_0 a \Gamma(b + 1)}{\rho q} \right]^{\frac{1}{b+1}}$$

The rate of change of the quantity q due to a particular microphysical process is given by:

$$\frac{\partial q}{\partial t} = \frac{1}{\rho} \int_0^{\infty} \frac{\partial m}{\partial t} N(D) dD$$

where $\frac{\partial m}{\partial t}$ is the rate of change of the mass of a single particle

Condensation-sublimation

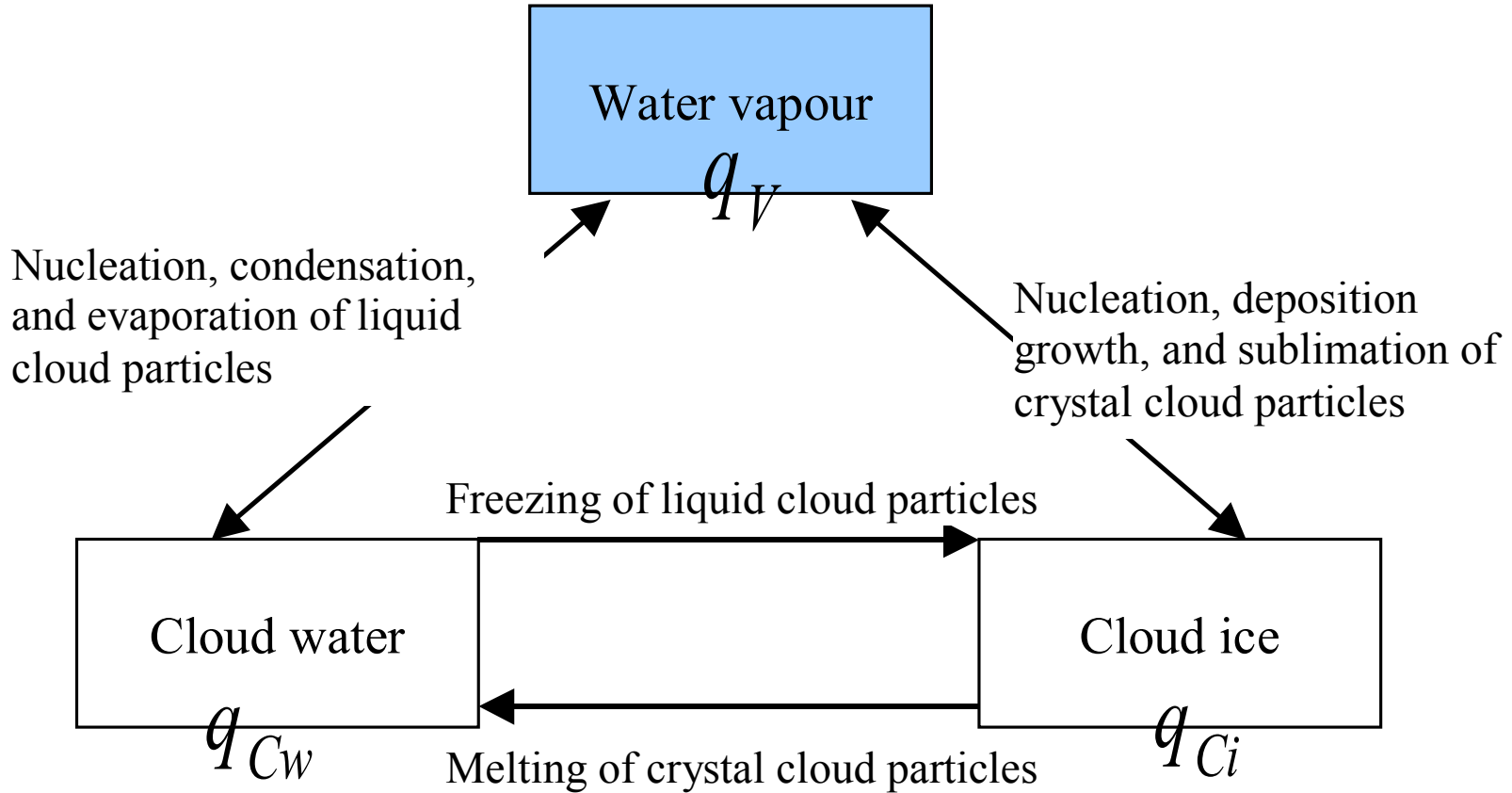
$$\frac{dm}{dt} = D \cdot \frac{2\pi F \left(\frac{q_v}{q_{sk}} - 1 \right) \rho}{\frac{1}{q_{sk} \chi} + \frac{L_k^v M_w}{K_a T} \left(\frac{L_k^v M_w}{R^i T} - 1 \right)} \cdot \left[1 - \frac{1}{2} \left(\frac{q_v}{q_{sk}} - 1 \right) \left[\frac{\rho \left(\frac{L_k^v M_w}{R^i T} - 1 \right)}{\frac{K_a T}{q_{sk} \chi L_k^v} + \rho \left(\frac{L_k^v M_w}{R^i T} - 1 \right)} \right]^2 \right] \left[1 + \frac{1 - 2 \frac{L_k^v M_w}{R^i T}}{\left(\frac{L_k^v M_w}{R^i T} - 1 \right)^2} \right]$$

where F is the ventilation coefficient, equal to 0.8 for cloud particles. For precipitation particles the following expression is implemented:

$$F = 0.78 + Sc^{1/3} \left(\frac{DU \rho}{\mu_{dif}} \right)^{1/3}$$

where μ_{dif} is the dynamical molecular viscosity of air, U the terminal velocity of the particle, and Sc the Schmidt number (= 0.6). The suffix k can be w or i , indicating liquid water or ice, respectively. L_w^v and L_i^v are the condensation and sublimation latent heat, χ the coefficient of molecular diffusion of vapour into air, K_a the thermal conductivity of air, M_w the molecular weight of water, and R^* the universal gas constant.

Fast microphysical processes



‘Autoconversion’

$$\frac{\partial q_{Cw,Ci}}{\partial t} = -q_{Cw,Ci} \frac{\Gamma(\alpha+b+1, \beta D_0)}{\Delta t \Gamma(\alpha+b+1)}$$

Fall of precipitation

The terminal velocity of one precipitation particle:

$$u(D) = kD^n \left(\frac{p_0}{p} \right)^{0.4}$$

where $n=0.8$ and $k=842 \text{ m}^{1-n}\text{s}^{-1}$ for rain and function of the type of ice particle for snow/hail

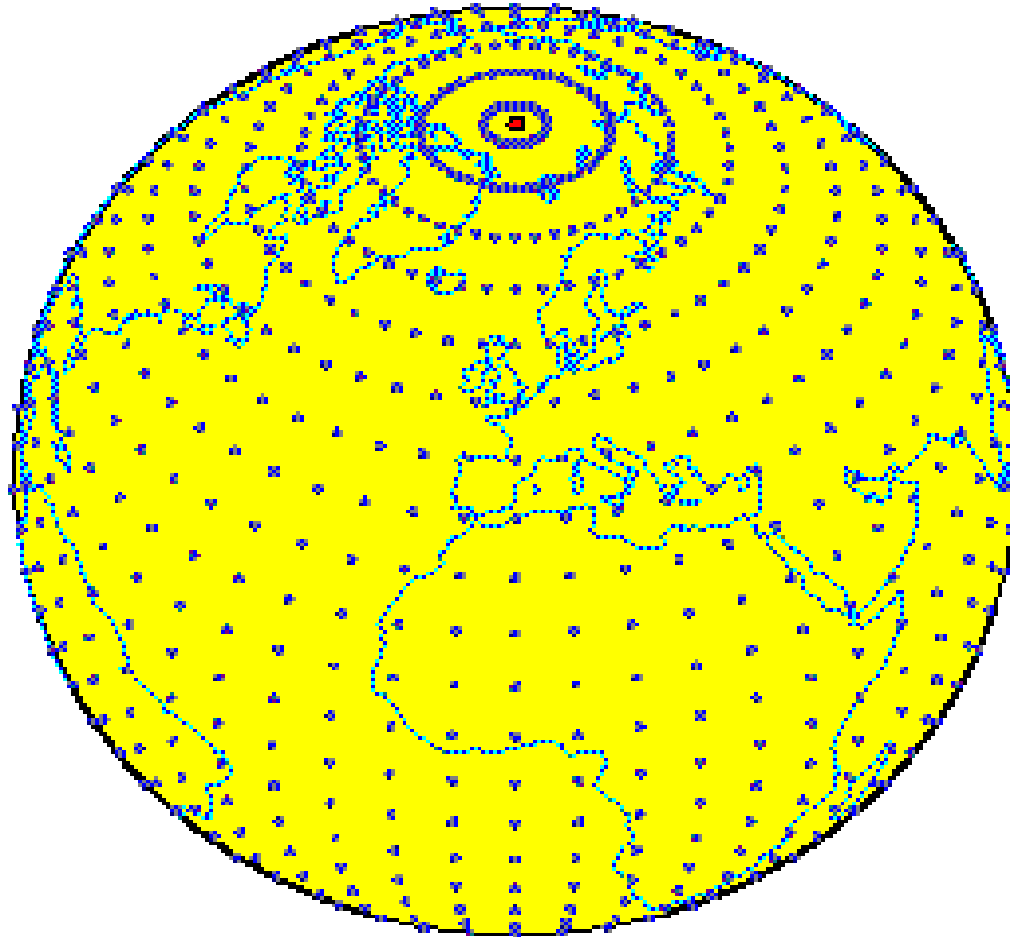
Averaged terminal velocity:

$$U = \frac{\int N(D)m(D)u(D)dD}{\int N(D)m(D)dD}$$

Conclusioni

'GLOBO'

GLOBAL version of the BOlam model



Model dynamics

- Hydrostatic, primitive equations
- Arakawa C grid; terrain-following, hybrid coordinates
- Explicit time split, Forward-Backward for gravity waves
- Advection: Weighted Average Flux (Toro 1989; Hubbard & Nikiforakis, 2001)
- Fourth order horizontal diffusion and second order divergence damping
- Polar filter (spectral along longitude)

Model physics:

- Radiation (Morcrette or Geleyn)
- Vertical diffusion (*E-l* scheme)
- Surface turbulent fluxes (Monin-Obuckov)
- Large scale precipitation and microphysics based on Shultz (1988)
- Moist convection based on the Kain-Fritsh parameterization
- Soil water and energy balance scheme based on Pressman (1994)
- Vegetation effects (Noilhan J., Mahfouf J.-F. ,1996)
- Gravity wave drag

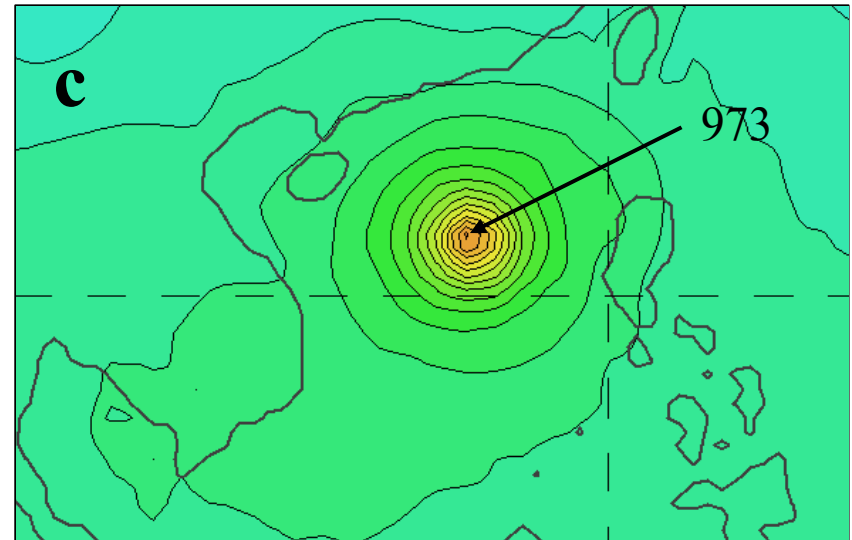
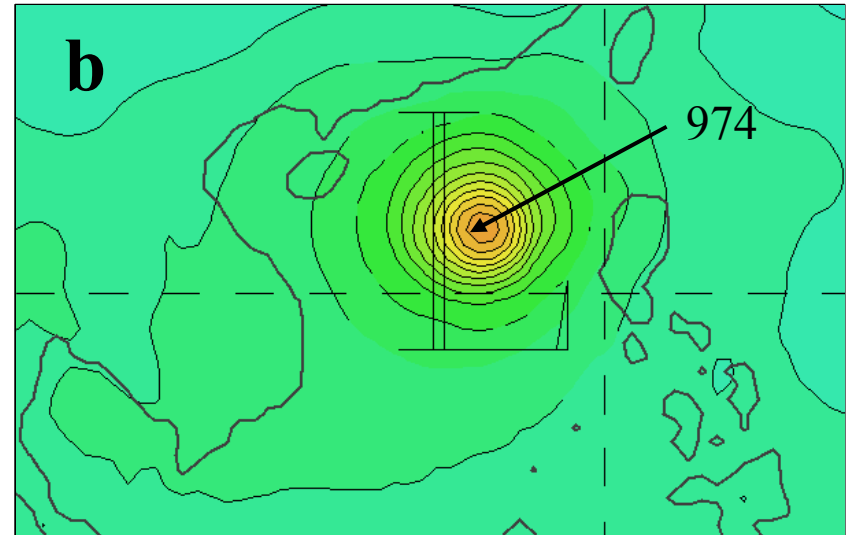
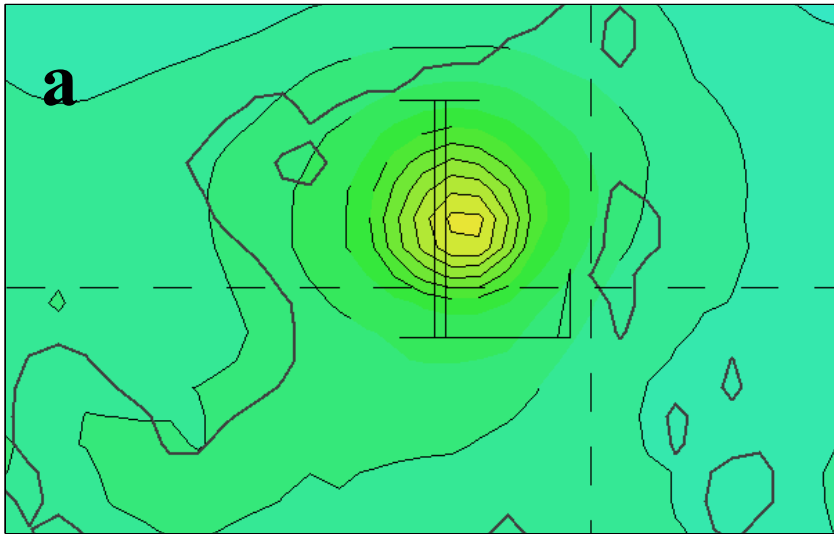
Case study : 2006/05/15

7-day forecast simulation, starting from ECMWF analysis at 00 UTC of May 15, 2006

3-D fields u , v , T , q extracted from *MARS* archive on 26 model levels and 1.0x1.0 lat-lon regular grid (12 Mbyte compressed).

2-D fields: soil temperature and water content (4 layers), snow height, log. of surface pressure, orography, and land sea mask.

System time: 1 hour (8 hours) at 1.0 (0.5) resolution on AMD64x2 with PG Fortran and MPI



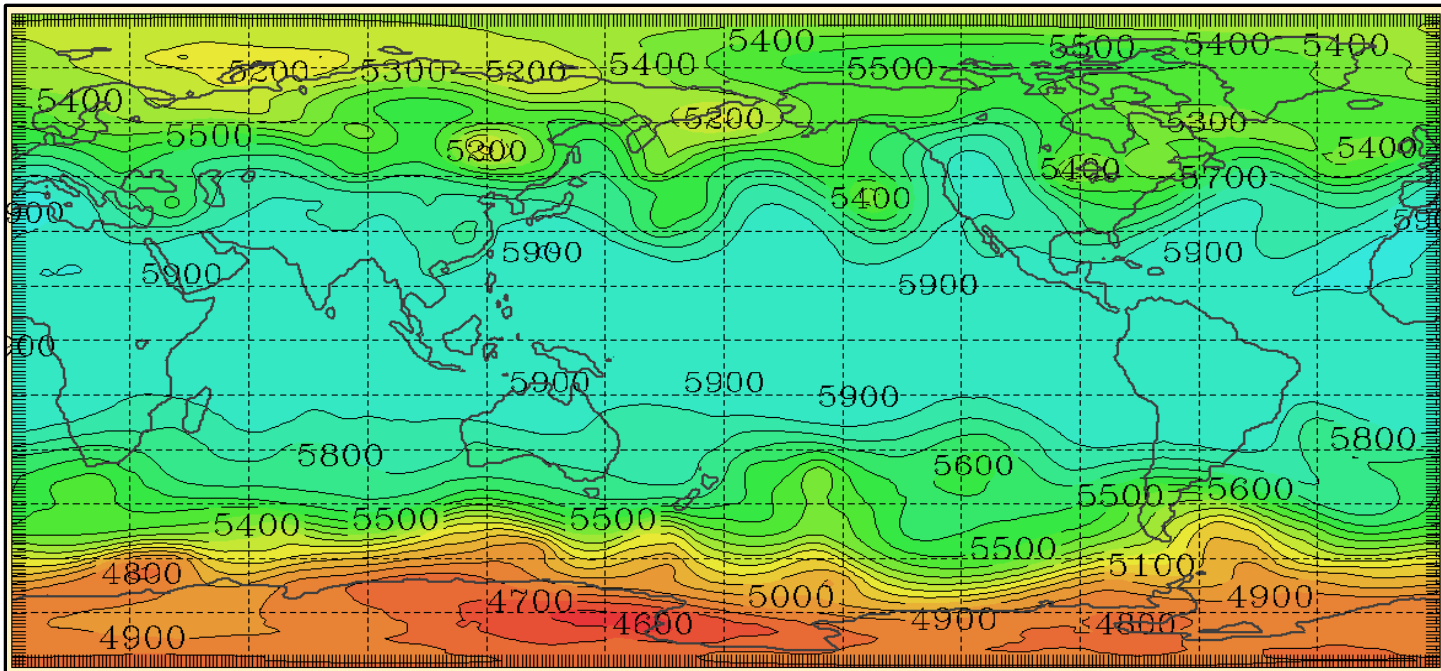
MSLP - 24 h forecast

- a) GLOBO 1.0 resolution
- b) GLOBO 0.5 resolution
- c) ECMWF (~ 0.25 resolution)

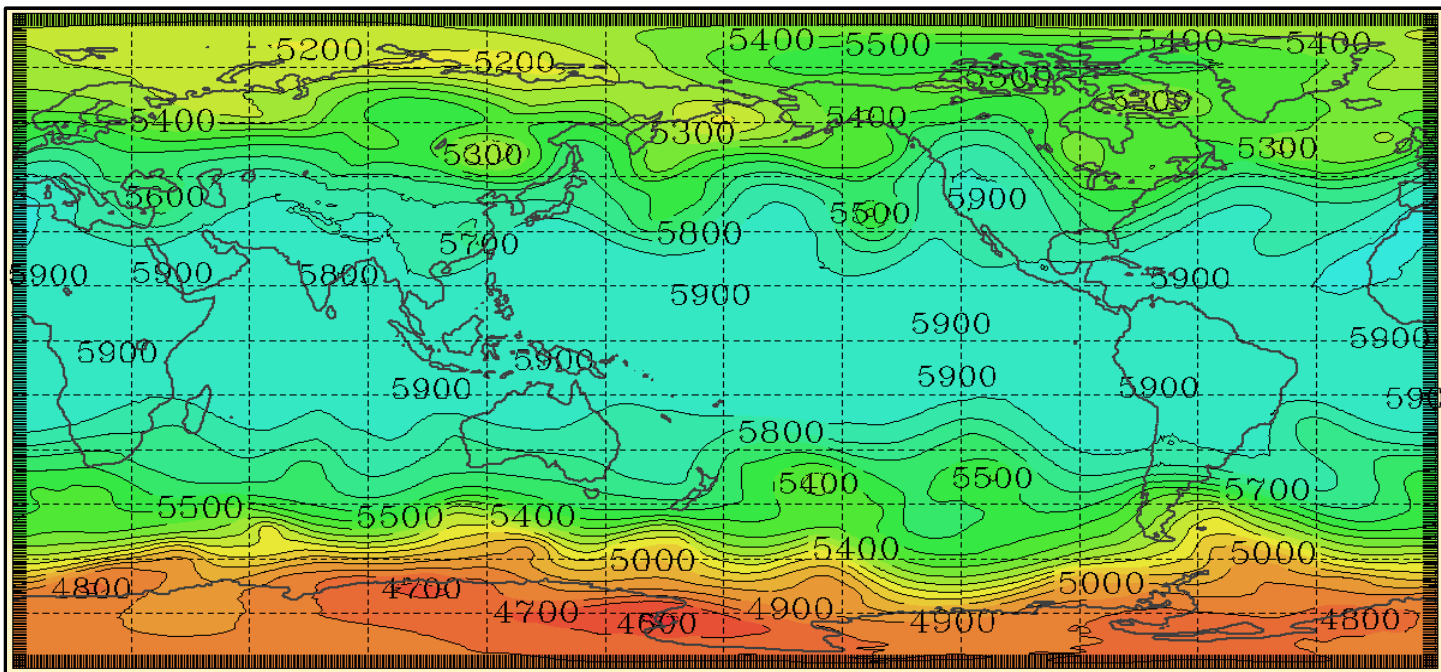
3-days Forecast

GLOBO vs ECMWF
500 hPa geopotential height

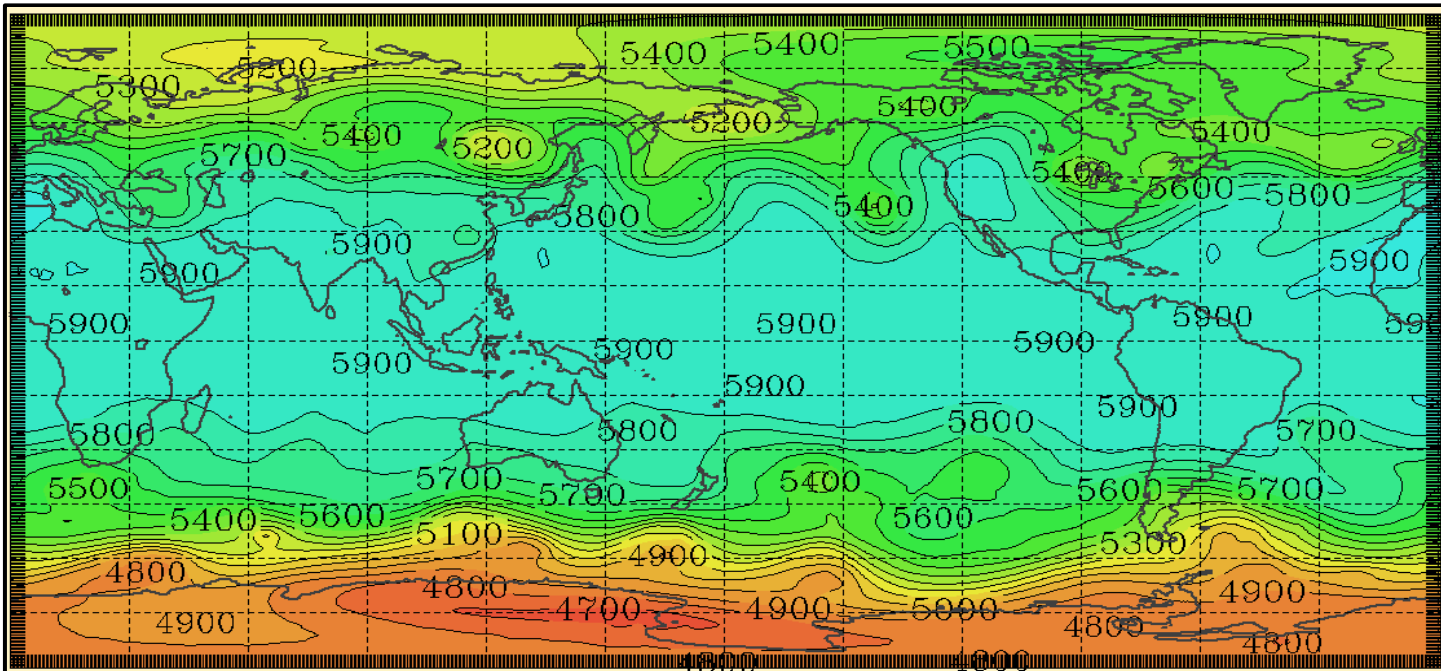
GLOBO 1.0x1.0 res.



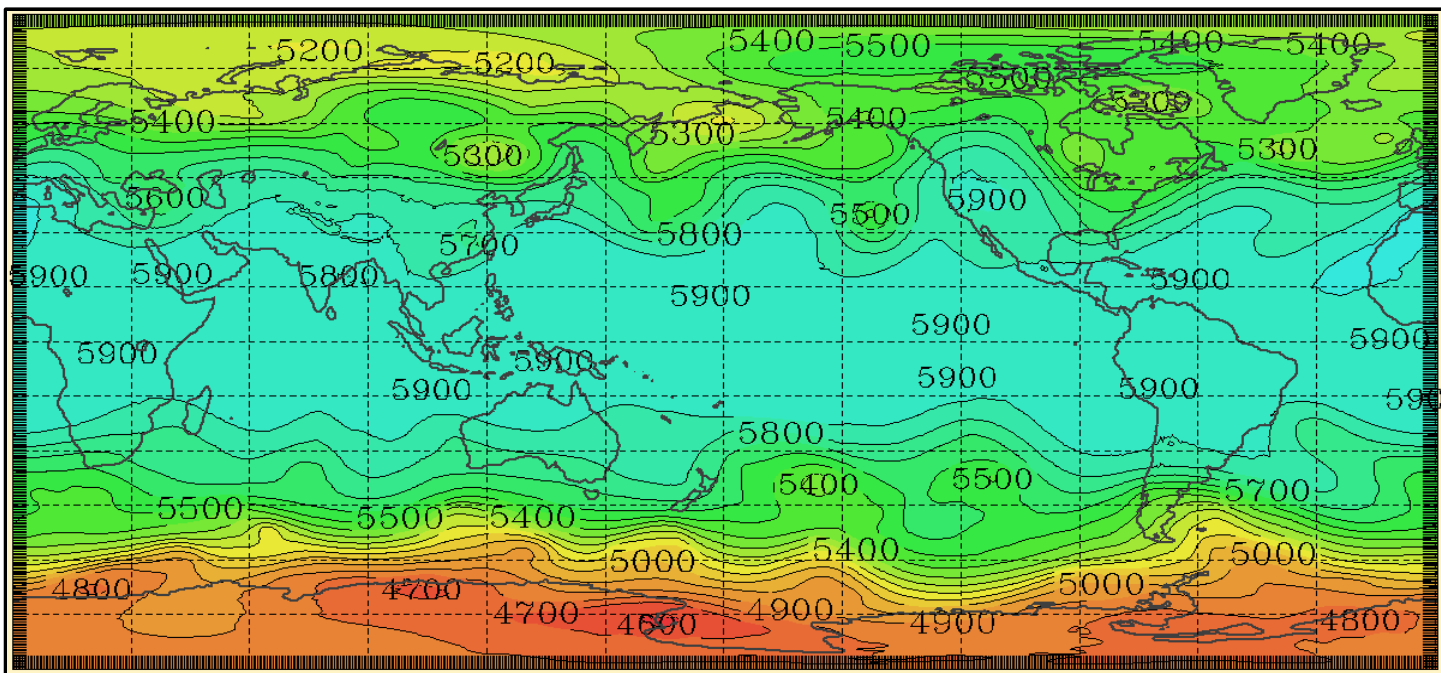
ECMWF



GLOBO 0.5x0.5 res.



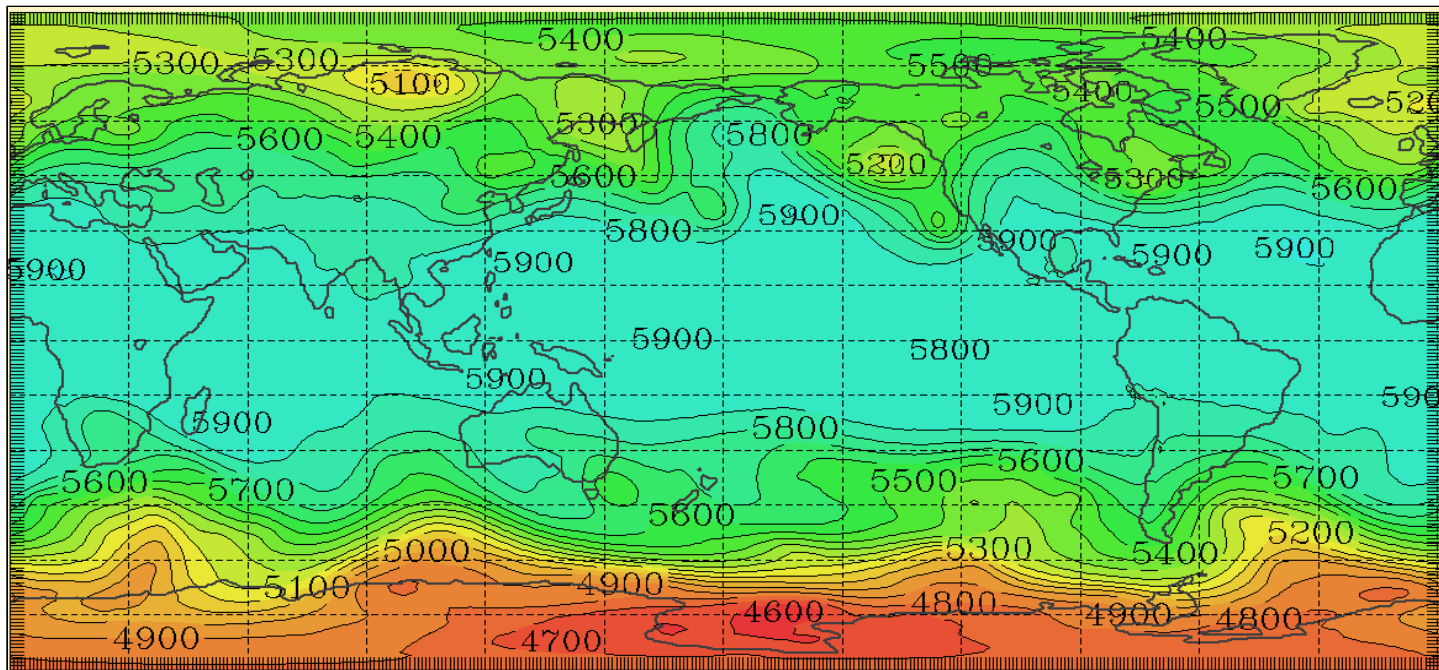
ECMWF



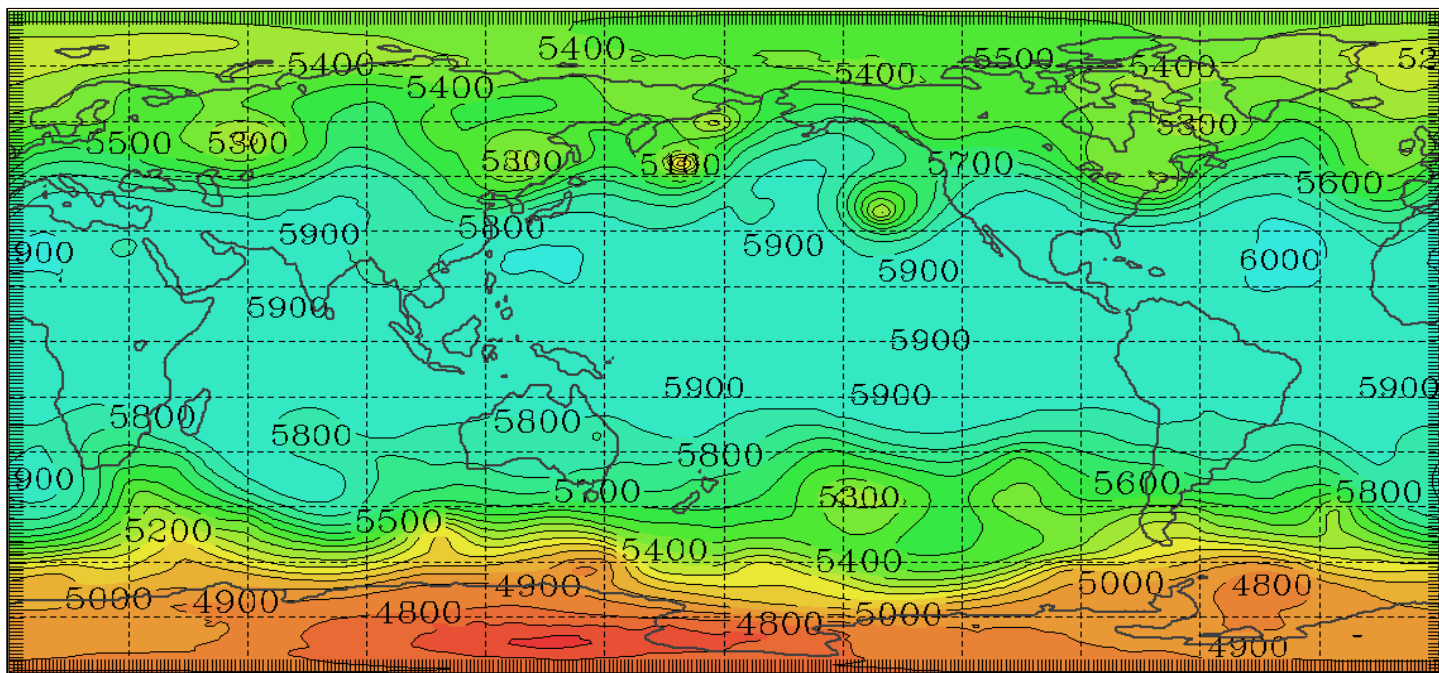
Medium range forecast (7 days)

**Comparison of Globo at 1.0 and deg resolution versus the
Ecmwf forecast and analysis**

Analysis

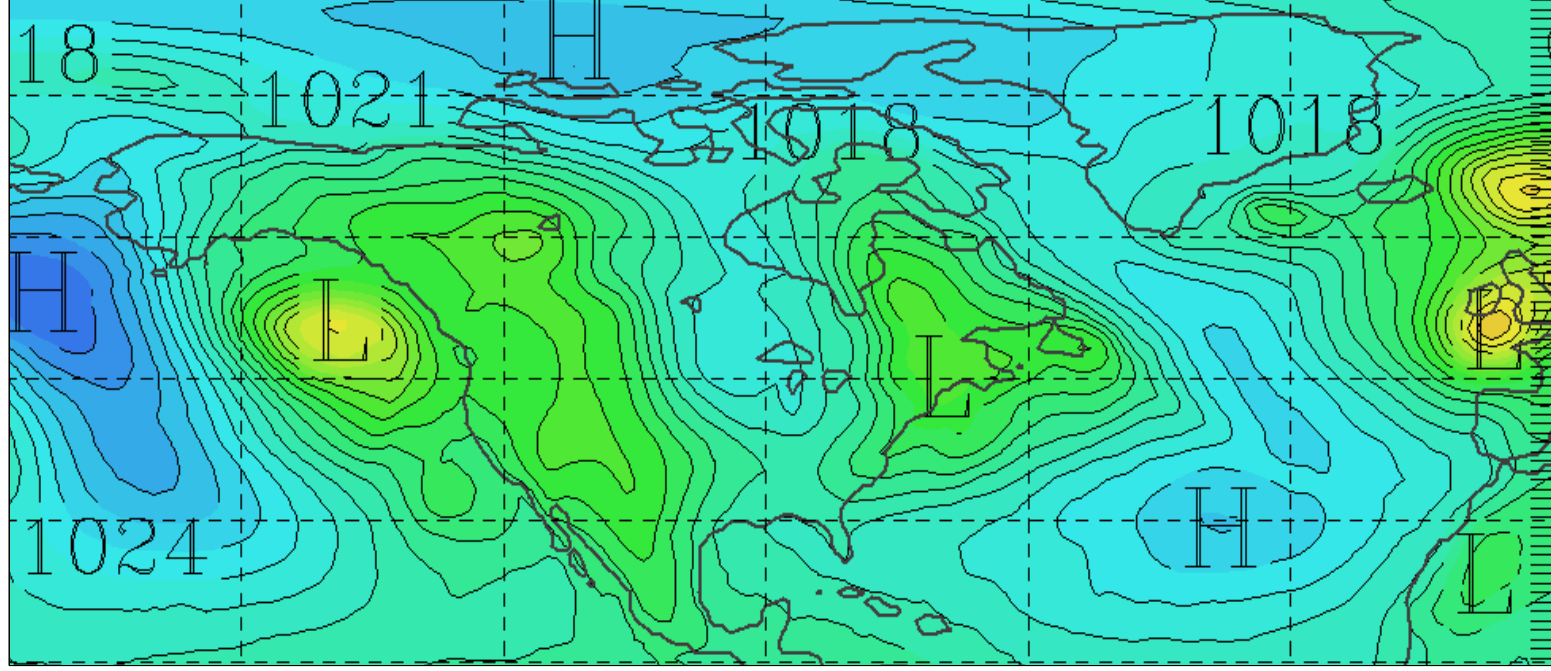
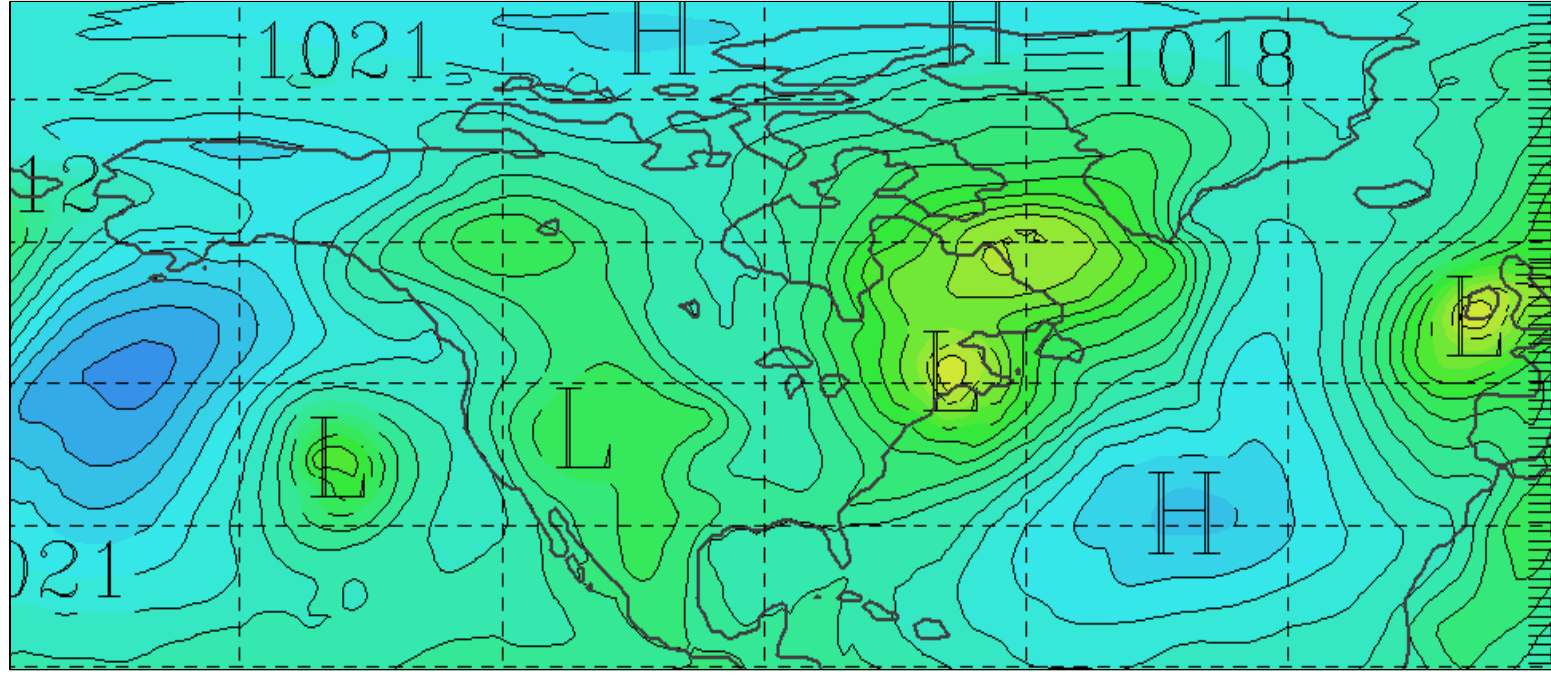


GLOBO 1.0x1.0 res.

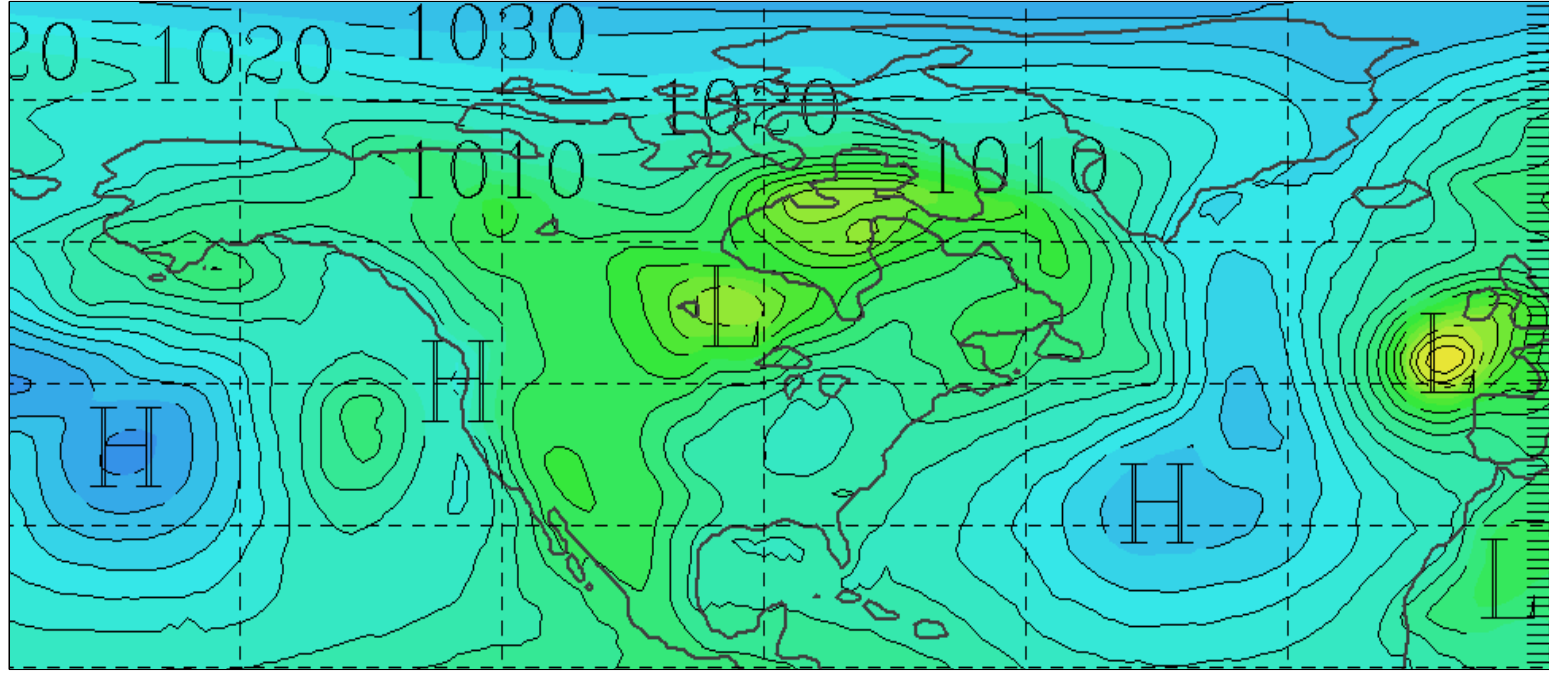


GLOBO 1.0x1.0 res.

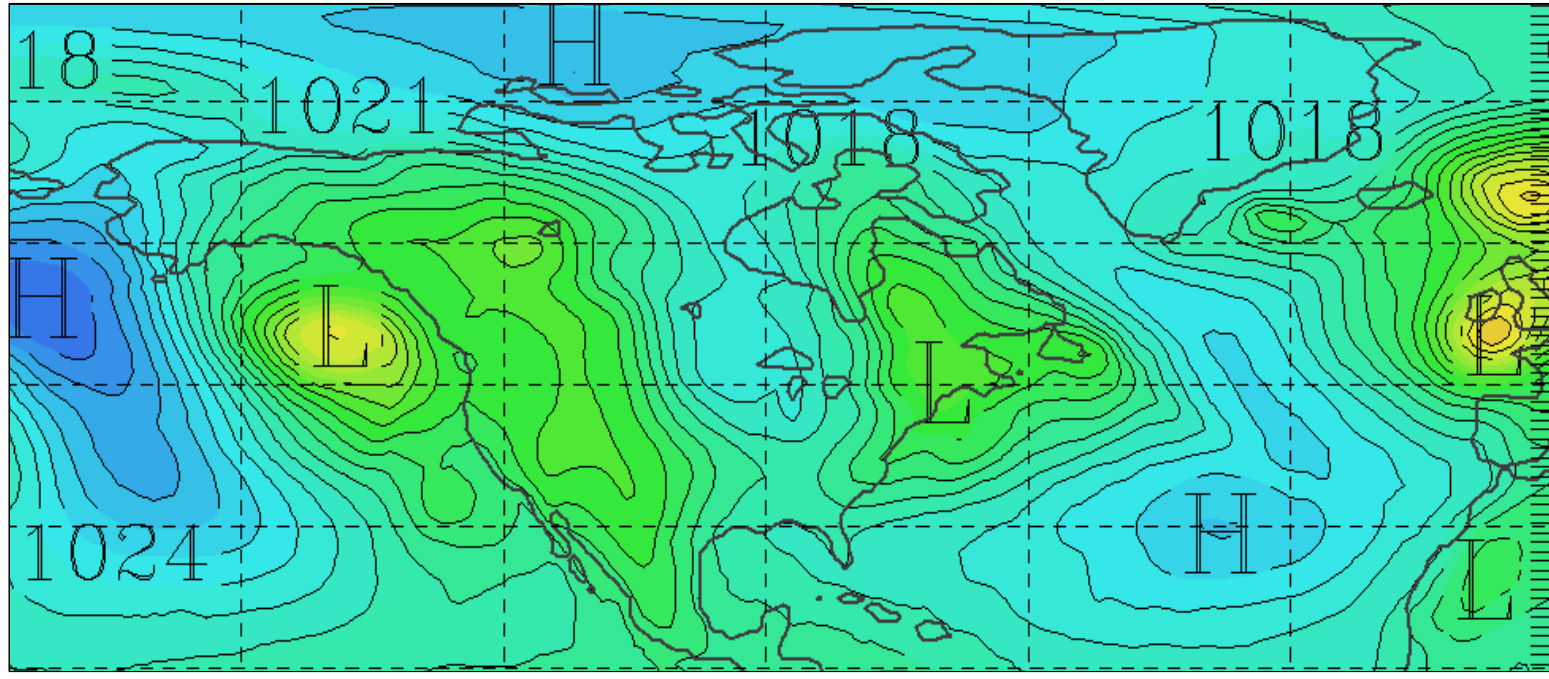
analysis 2006/05/22



ECMWF



analysis 2006/05/22



Objectives

- Develop a tool for medium range weather and ensemble forecast, to be used by weather services;
- AGCM to be used to study the general circulation of the atmosphere (role of water vapour, baroclinic instability, low-frequency variability, planetary waves,)
- Tool for seasonal prediction (?)
- **Climate model ?**